Contrasting responses to drought of forest floor CO\textsubscript{2} efflux in a Loblolly pine plantation and a nearby Oak-Hickory forest

S. PALMROTH\textsuperscript{*}, C. A. MAIER\textsuperscript{t}, H. R. McCARTHY\textsuperscript{*}, A. C. OISHI\textsuperscript{*}, H.-S. KIM\textsuperscript{*}, K. H. JOHNSON\textsuperscript{j}, G. G. KATUL\textsuperscript{*} and R. OREN\textsuperscript{*}

\textsuperscript{*}Nicholas School of the Environment and Earth Sciences, Duke University, Box 90328, Durham, NC 27708-0328, USA, \textsuperscript{t}Department of Forest Ecology, University of Helsinki, PO Box 27, 00014 Helsinki, Finland, \textsuperscript{j}Southern Research Station, USDA Forest Service, 3041 Coralvills Road, Research Triangle Park, NC 27709, USA

Abstract

Forest floor CO\textsubscript{2} efflux ($F_{HF}$) depends on vegetation type, climate, and soil physical properties. We assessed the effects of biological factors on $F_{HF}$ by comparing a maturing pine plantation (PP) and a nearby mature Oak-Hickory-type hardwood forest (HW). $F_{HF}$ was measured continuously with soil chambers connected to an IRGA during 2001–2002. At both sites, $F_{HF}$ depended on soil temperature at 5 cm ($T_{5}$) when soil was moist (soil moisture, $\theta > 0.20$ m\textsuperscript{3} m\textsuperscript{-3}), and on both $T_{5}$ and $\theta$ when soil was drier. A model ($F_{HF} (T_{5}, \theta)$) explained $\geq$ 92\% of the variation in the daily mean $F_{HF}$ at both sites. Higher radiation reaching the ground during the leafless period, and a thinner litter layer because of faster decomposition, probably caused higher soil temperature at HW compared with PP. The annual $F_{HF}$ was estimated at 1330 and 1464 g C m\textsuperscript{-2} yr\textsuperscript{-1} for a year with mild drought (2001) at PP and HW, respectively, and 1231 and 1557 g C m\textsuperscript{-2} yr\textsuperscript{-1} for a year with severe drought (2002). In the wetter year, higher soil temperature and moisture at HW compared with PP compensated for the negative effect on $F_{HF}$ of the response to these variables resulting in similar annual $F_{HF}$ at both stands. In the drier year, however, the response to soil temperature and moisture was more similar at the two stands causing the difference in the state variables to impel a higher $F_{HF}$ at HW. A simple mass balance indicated that in the wetter year, C in the litter–soil system was at steady state at HW, and was accruing at PP. However, HW was probably losing C from the mineral soil during the severe drought year of 2002, while PP was accumulating C at a lower rate because of a lack of C from the litter layer. Such contrasting behavior of two forest types in close proximity might frustrate attempts to estimate regional carbon (C) fluxes and net C exchange.

Keywords: belowground carbon allocation, chamber, conifer, hardwood, IRGA, litterfall, Pinus taeda L., Quercus ssp., soil respiration, temperate

Received 27 May 2004; received in revised form 10 September 2004; accepted 20 September

Introduction

The importance of soils in global carbon (C) accounting requires that soil C fluxes and stocks be accurately quantified across ecosystems and climates (Raich & Nadelhoffer, 1989; Raich & Schlesinger, 1992; Liski et al., 2003; Reichstein et al., 2003). At the ecosystem level, net ecosystem exchange (NEE) of carbon (C) is defined as the difference between gross primary production (GPP), the amount of C assimilated in photosynthesis, and ecosystem respiration ($R_{E}$), the amount of C re-emitted to the atmosphere from autotrophic and heterotrophic respiration. $R_{E}$ utilizes a large proportion of GPP, with an average ratio for $R_{E}$/GPP of 0.82 for temperate coniferous forests and 0.77 for temperate broad-leaved deciduous forests (FluxNet sites; Falge et al., 2002). Forest floor CO\textsubscript{2} efflux ($F_{HF}$), in turn, represents a large proportion of $R_{E}$ with an average $F_{HF}$/ $R_{E}$ ratio of 0.69 (EuroFlux sites; Janssens et al., 2001).
Consequently, accurate predictions of NEE responses to variation in climate depend on understanding and quantifying the responses of \( F_h \) to climatic variables in different forest types.

Climate, net primary production (NPP), and \( F_h \) are linked through the closely coupled C and nutrient cycles (Raich & Schlesinger, 1992). In most cases, warmer and wetter climates support higher productivity and, thus, higher litterfall and absolute amount of belowground C allocation. At the global scale, \( F_h \) has been related to NPP and temperature (Raich & Schlesinger, 1992). However, other studies at large geographical scales have found that \( F_h \) is unrelated to temperature (Janssens et al., 2001; Reichstein et al., 2003). In these studies \( F_h \) was related to precipitation and productivity indices, such as leaf area index (LAI) and leaf production. The prevalence of studies that find a productivity measure to explain a large proportion of the variation in \( F_h \) is not surprising. According to a simple mass balance, the sum of litterfall and belowground allocation equals \( F_h \) if changes in soil and litter C stocks, and C transported off site are assumed negligible relative to the other fluxes (Raich & Nadelhoffer, 1989; Giardina & Ryan, 2002). A recent study by Davidson et al. (2002) suggested that, when averaged across biomes and stands of different ages, the annual release of C in \( F_h \) is roughly four times the amount of C in the annual aboveground litterfall.

Forest floor \( \text{CO}_2 \) efflux reflects both the production of \( \text{CO}_2 \) and its transport to the surface of the litter layer, although integrated over long periods \( F_h \) reflects primarily the production of \( \text{CO}_2 \) in respiration. The production of \( \text{CO}_2 \) is comprised of root and fungal respiration, and microbial decomposition of dead roots, root exudates, fungal hyphae, and leaf and woody litter (e.g., Hanson et al., 2000). Thus, the production processes must be affected by vegetation type, which determines the quantity, quality, and timing of litter fall, root biomass and its turnover rate, and photosynthetic activity and allocation patterns of recent photosynthates (Hanson et al., 2000; Hogberg et al., 2001). The production of \( \text{CO}_2 \) in the soil creates a concentration gradient driving the gaseous transport, which is affected by the physical properties, temperature, and the water content of the soil–litter continuum (e.g., Nobel, 1999). Thus, temperature and precipitation greatly influence, directly or indirectly, all the component processes of \( F_h \) (e.g., Singh & Gupta, 1977; Raich & Shlesinger, 1992; Lloyd & Taylor, 1994; Reichstein et al., 2003).

Vegetation cover affects local soil temperature through intercepted radiation and local soil moisture through interception and transpiration. Although Raich & Tufekcioglu (2000) suggested that soil respiration is controlled primarily by climate factors with vegetation only having a secondary effect, soil respiration rates in coniferous forests were 10% lower, on average, than those in adjacent broad-leaved stands on the same soil type. The reasons underlying this finding remain largely unresolved (Raich & Tufekcioglu, 2000), and may relate to the effect of vegetation on the seasonal dynamics of soil respiration (Curiel Yuste et al., 2004). The difference in magnitude of soil respiration between different forest types may have a large effect on NEE of the forests, particularly in regions with warm, moist temperate climate that support high rates of productivity and \( F_h \) such as the south-east US.

The forested area of the south-eastern US is a mosaic composed largely of pine forest (both naturally and artificially regenerated), mixed pine–hardwood forests, and hardwood forests (HW). Pine forests cover ~0.13 million km\(^2\) (Wear, 2002) and can support very high GPP, NPP, NEE, and \( F_h \) (Clark et al., 1999; Moncrieff & Fang, 1999; Maier & Kress, 2000; Andrews & Schlesinger, 2001; Katul et al., 2001; Finzi et al., 2002; Butnor et al., 2003; Schäfer et al., 2003). Relative to pine forests, less has been published on this region’s hardwood forests (Hanson et al., 1993; Curtis et al., 2001; Norby et al., 2002; Lee & Jose, 2003; King et al., 2004), which cover ~0.30 million km\(^2\). However, the area under pine is projected to increase by 0.09 million km\(^2\) by 2040 while the area under HW is projected to decrease by 0.11 million km\(^2\) (Wear, 2002). Thus, our ability to quantify differences in \( F_h \) between these forest types is critical for estimating current and future C budgets for the region.

The aim of this study was to quantify seasonal and annual \( F_h \) in both forest types, and to separate the effect of forest type from that of climate or soil on the magnitude and dynamics of \( F_h \). To accomplish this we compared \( \text{CO}_2 \) efflux in a maturing loblolly pine plantation (PP) and a mature Oak-Hickory type forest (HW), representative of a large portion of south-eastern US forests, during a mild and severe drought years (2001 and 2002, respectively). The two experimental stands represent different developmental stages. Maximum tree height at HW was twice that at PP, and aboveground living biomass ~27% higher at HW; however, in 2001, peak one-sided LAI and litterfall were more similar at the two stands (Table 1). We present a unique combination of synchronous and near-continuous monitoring of \( F_h \) at both HW and PP (in temporal scales from ~3 h to 2 years), large spatial coverage by the monitoring system, and close proximity of the two study stands (~1 km apart). HW and PP were, therefore, on a similar soil type, and experienced identical incident radiation and precipitation above the canopy. Hence, differences between the stands in the key forcing variables, such as soil temperature and moisture, are solely reflecting differences in the amount of leaf area and its dynamics, transpiration, and litter quality.
Table 1  Maximum height, total aboveground living biomass (in September 2002), peak one-sided leaf area index (LAI), and annual litterfall at the hardwood forest (HW) and the pine plantation (PP) located at Duke Forest

<table>
<thead>
<tr>
<th></th>
<th>Year</th>
<th>HW</th>
<th>PP</th>
</tr>
</thead>
<tbody>
<tr>
<td>Height (m)*</td>
<td>2002</td>
<td>35</td>
<td>17</td>
</tr>
<tr>
<td>Total aboveground living biomass (kg m(^{-2}))</td>
<td>2002</td>
<td>14.1 (1.23)</td>
<td>10.8 (0.28)</td>
</tr>
<tr>
<td>Peak LAI</td>
<td>2001</td>
<td>6.9 (0.25)</td>
<td>5.6 (0.25)</td>
</tr>
<tr>
<td>Litterfall (g C m(^{-2}))</td>
<td>2002</td>
<td>342 (36)</td>
<td>332 (16)</td>
</tr>
</tbody>
</table>

Standard error in parenthesis; N = 3 at HW and 5 at PP.
*In the beginning of 2002 at PP.
†Biomass equations from Brown et al. (1997) and Naidu et al. (1998).

Materials and methods

Site description

The PP, planted in 1983, and the 80-100-year-old uneven aged HW, are both AmeriFlux sites located at the Duke Forest C-H\(_2\)O Research Site, in Orange County, NC (35°58'N, 79°08'W). The soil is classified as Enon silt loam, a low fertility Hapludalf typical of the SE US Piedmont, with a transition to Iredell gravelly loam toward HW (Pataki & Oren, 2003). An impervious clay pan underlies the research sites at ca. 30 cm belowground. The topographic variations are small (<5% slopes). The region is characterized by warm summers and mild winters. The annual mean temperature is 15.5°C and the 1140 mm of annual precipitation is evenly distributed throughout the year.

At PP, Pinus taeda L. is in the dominant canopy position together with fewer individuals of Liquidambar styraciflua L. The subcanopy contains ~40 woody species of which Acer rubrum L., Ulmus alata Michx., and Cornus florida L. are the most prevalent. The HW stand is composed mostly of Liriodendron tulipifera L., Quercus alba L., Q. michauxii Nutt., Q. phellos L., L. styraciflua L., and Carya ssp. The stand characteristics are given in Table 1.

Forest floor CO\(_2\) efflux and litterfall

\(F_d\) was measured with the automated carbon efflux system (ACES, US Patent 6692970) developed at USDA Forest Service, Southern Research Station Laboratory in Research Triangle Park, NC (Butnor et al., 2003). ACES is an open system where an IRGA is connected with 15 soil chambers (491 cm\(^3\)) equipped with air and soil (at 5 cm, \(T_s\)) thermocouples, pressure equilibration ports, and reflective covers. Forest floor CO\(_2\) efflux was measured for 10 min from each chamber. ACES gives consistent responses regardless of differences in soil and litter properties and has been calibrated with known CO\(_2\) efflux rates (Butnor & Johnsen, 2004).

An ACES unit with eight chambers was operating at PP from February 2001 until August 2001 and with six chambers until December 2002. At HW, a unit with 15 chambers was operating from June 2001 until December 2002. Chambers were placed randomly at PP and systematically at HW because of coordination with eddy covariance measurements at HW. To minimize the effect of precipitation and litterfall exclusion on the soil substrate within the chambers, they were moved every 3-4 days between two sample points (i.e., there were two locations, A and B, for each chamber). Metal collars attached to the chambers (and a collar only for the ‘empty’ location) were pushed 1-2 cm into the mineral soil. To minimize long-term disturbance caused by the experiment, the chambers at both stands were relocated on March 26-28, 2002. Starting in the fall of 2002, litter from litter collectors was used to replace the litter excluded while the chambers were in place. The litter was weighed in the field (wet weight), and distributed to each A and B location in proportion to the area covered by the chamber.

Litterfall collection began at both stands in September 2001. There were 20 traps (0.5 m\(^2\)) at PP and 48 at HW. These were emptied weekly or biweekly during the heaviest litterfall period and monthly or bimonthly otherwise. Leaf area of a subsample was measured with a leaf area meter (DIAS, Decagon Devices Inc., Pullman, WA, USA) and all the material (divided into leaves, woody material, seeds, and other) was dried (≥ 3 days at 68°C) and weighed. C content of 0.47 was used in dry-mass-to-C conversions (Hamilton et al., 2001). In 2001, the annual litterfall at PP was extrapolated using data from the free-air CO\(_2\) enrichment (FACE) experiment (Finzi et al., 2002), and at HW by assuming that monthly litterfall, excluding leaves, is constant throughout the year. At HW, LAI was derived from plant area index measurements with LAI-2000 (Li-Cor Inc., Lincoln, NE, USA) scaled with the estimates of LAI obtained from the litterfall data. At PP, it was calculated as the balance between growth of new foliage (needle elongation measurements) and litterfall.

Environmental data

Environmental data were partially obtained from the nearby FACE experiment (Duke Forest FACE (FACTS I); http://face.env.duke.edu). Precipitation (mm) was measured with one tipping bucket (TI, Texas Instruments, Austin, TX, USA) above the canopy. At each of the six FACE plots at PP, volumetric soil moisture
content (θ, m$^3$ m$^{-3}$) was measured with four frequency-domain reflectometry probes (CS615 Campbell Scientific, Logan, UT, USA) placed in the upper 30 cm of the mineral soil. Soil temperature ($T_{soil}$, °C) was measured with one sensor (M 841/51, Siemens, Germany) at 10–15 cm depth at each FACE plot. At HW, θ was measured with theta probes (ML1X/ML2X, Delta-T Devices, Cambridge, UK) placed at 10 cm (two probes) and 25 cm (two probes) in the mineral soil. The site mean θ for HW was obtained by averaging over the four ML1X/ML2X probes, and at PP by averaging over all CS615 probes located in the ambient FACE plots. Comparison of the theta probe measurements at HW with identical θ probe measurements at PP showed similar inter-site difference in θ to that suggested by the data used in this study (H. McCarthy, unpublished data). All sensors were sampled every 30 s, and data were averaged over 30 min and stored in a data logger (21X or CR23X, Campbell Scientific).

Data analysis

The response of $F_{H}$ (μmol CO$_2$ m$^{-2}$ s$^{-1}$) to temperature ($T_{soil}$, °C) and volumetric soil moisture content (θ, m$^3$ m$^{-3}$) was described as

$$F_{H}(T_{soil}, \theta) = R_0 e^{a T_{soil}} [1 - e^{-(b\theta + c)}],$$

where $R_0$ (μmol m$^{-2}$ s$^{-1}$) is the intercept at 0 °C, the so-called base respiration, and $a$ is the temperature sensitivity ($Q_{10} = e^{a T_{soil}}$) when soil moisture is not limiting, representing potential efflux $F_{H}^{P}$. The function $1 - e^{-(b\theta + c)}$, where $b$ and $c$ are fitting parameters, gives the reduction of $F_{H}$ from $F_{H}^{P}$ as soil moisture decreases (modified from Fang & Moncrieff, 1999). The threshold for 'nonlimiting soil moisture' was defined as the value of θ where $F_{H}$ reached 90% of the $F_{H}$ averaged over the range of θ within which $F_{H}$ was clearly insensitive to soil moisture ($θ > 0.25$ m$^3$ m$^{-3}$). Very high values of soil moisture ($θ > 0.4$) occurred for short periods only and no reduction in $F_{H}$ (because of oxygen limitation) was observed at either stand. The curve fitting was done using the nonlinear curve fitting procedure of Systat (Systat Software Inc., Richmond, CA, USA).

To study the relationship between $R_0$, $a$, and θ (i.e., the assumption that these parameters are independent of θ) the data were filtered for $T_{soil}$ values ranging from 12 to 22 °C at PP, and from 12 to 24 °C at HW, and $θ < 0.3$ m$^3$ m$^{-3}$ at both stands. Consequently, the conditions containing very few observations (low $T_{soil}$ and θ, and high $T_{soil}$ and θ) were excluded. The remaining data were stratified into soil moisture classes (10 at PP, and nine at HW) with approximately equal number of observations in each class, and the parameter values were estimated for each bin. To study the temperature dependence of the parameters of the $F_{H}$-temperature response, the data were filtered for nonlimiting soil moisture. $R_0$ and $a$ were then estimated for five temperature bins at PP and six bins at HW. Although each chamber location (i.e., location A or B for each chamber) was considered a replicate, treating those separately would have resulted in too few observations within each class. Therefore, we used instantaneous data pooled over all the chamber locations (pooled data) in this analysis.

To test for differences between stands in the $F_{H}$-temperature response, we quantified the within-stand spatial variability by estimating the parameters separately for each chamber position (location data). The number of chambers multiplied by two locations for each chamber allowed the generation of 12-16 and 30 relationships (replicates) at PP and HW, respectively. The normality of the distribution of the values of $R_0$ and $a$ (Eqn 1) was tested using the Lilliefors modification of the Kolmogorov–Smirnov test (Systat Software Inc.). The differences between stands in the mean values of $R_0$ and $a$ were studied using ANOVA (Systat Software Inc.). The values of $R_0$ and $a$, averaged for each stand, were used to calculate daily mean $F_{H}^{P}$. The daily mean of the measured $F_{H}$ (daily data) was divided by the respective daily $F_{H}^{P}$ and this ratio was used in estimation of the parameter values of the $F_{H}$-soil moisture reduction function. The between-stand difference in the parameter values of this response was tested by comparing a full model (with stand-specific parameters) to a reduced model (with a single set of parameters) based on F-test statistics for extra sum of squares (Ramsey & Schaefer, 1997).

Daily averaging did not result in a considerable loss of information or introduction of bias because of the nonlinearity of the response functions. At both stands, the average daily range (max–min) in $T_{soil}$ in 2002 was ~ 2 °C, and >90% of the observed ranges were < 3.0 °C. The average daily range in θ was ~ 0.016 m$^3$ m$^{-3}$, and >93% of the observed ranges were < 0.05 m$^3$ m$^{-3}$. A daily mean datum was included in the analysis if ≥ 50% of the potential observations for the 24 h period were available. The number of days fulfilling this criterion was 467 at PP, and 409 at HW, and the proportion of gaps (in $F_{H}$ and $T_{soil}$ data) was 36% at PP, and 44% at HW over the study period in 2001–2002. Soil temperature for the entire study period was available at ~ 10 cm ($T_{soil}$) at the FACE site. In gap-filling, to avoid introducing an additional source of uncertainty when deriving $T_{soil}$ from $T_{soil}$, the model was reparameterized using $T_{soil}$. The relationships between $F_{H}$ and $T_{soil}$, and $F_{H}^{P}$ and $T_{soil}$ were equally good at both stands. The data were gap-filled (using Eqn 1) and monthly and annual estimates of $F_{H}$ for the two stands were compared.

© 2005 Blackwell Publishing Ltd, Global Change Biology 11, 421–434
Results

Daily mean soil temperature at 5 cm ($T_s$), volumetric soil moisture content ($\theta$), and forest floor CO$_2$ efflux ($F_{HF}$) were higher at HW compared with PP most of the time over the study period (2001–2002) (Figs 1a, b and d). Precipitation was evenly distributed throughout the 2 years except for a drought period in the last quarter of 2001 and the second quarter of 2002. Peak one-sided LAI and annual litterfall were higher at HW than at PP (Fig. 1c, Table 1), and the differences increased with increasing drought severity in 2002.

Variation in $F_{HF}$-temperature response

The multiplicative model (Eqn (1)) assumes that the parameters of the $F_{HF}$-temperature response function ($R_0$ and $a$) are independent of soil moisture. To test this assumption, the pooled data were stratified into soil moisture classes. Within the bins, the variation in $T_s$ explained more than 50% of the variation in $F_{HF}$ ($P<0.001$) in all but the driest bin ($r^2 = 0.24$). Although soil moisture appeared to have influenced $R_0$ and $a$, the effect was not clear and restricted to very low soil moisture ($\theta < 0.15 \text{ m}^3 \text{ m}^{-3}$; Figs 2a and c), thus, giving support to our modeling approach.

To study the temperature dependence of the parameters of the $F_{HF}$-temperature response, the pooled data were filtered for nonlimiting soil moisture (representing $F_{HF}^{\theta}$). $F_{HF}$ was considered as unlimited by soil moisture at $\theta \geq 0.2 \text{ m}^3 \text{ m}^{-3}$. The relationship between $F_{HF}$ and $T_s$ was significant in all the temperature bins ($P<0.001$; $N>1000$), but the explained variation in $F_{HF}$ was low ($\leq 24\%$) reflecting high 'noise-to-signal ratio' over the narrow range in $T_s$ within each bin (3°C). The values of $R_0$ increased and those of $a$ decreased with increasing temperature (Figs 2b and d). The inverse relationship between the two parameters obtained from Figs 2b and d is shown in Fig. 3a.

To quantify the spatial variation of the parameters of the $F_{HF}$-temperature response, $R_0$ and $a$ were estimated for each chamber location (i.e., using location data). In addition, based on a subsequent analysis of residuals, we partitioned the data into two periods, representing conditions before and after drought-breaking rains in 2002. From August 27, $\theta$ remained $\geq 0.20 \text{ m}^3 \text{ m}^{-3}$
for most of the year after staying below that value for ~120 consecutive days. The distribution of the sampled values of $R_b$ and $a$ at both stands and periods were normal ($P > 0.10$). A two-way ANOVA with stand and period as main effects was used to test for differences in $R_b$ and $a$. The time interval before and after the relocation of the chambers in March 2002 was incorporated as a nested effect in the period main effect, and was found insignificant ($P > 0.20$). The effects of both stand and period, and their interaction effect were significant ($P < 0.01$). The mean value of $R_b$ was lower and of $a$ higher at HW compared with PP during the first period (Tukey’s pairwise comparison, $P < 0.05$; Table 2).

An inverse relationship, similar to that found in analysis of the temperature bins (Fig. 3a), emerged when $R_b$ and $a$ were estimated for each chamber location (Fig. 3b). The comparison of a full model (stand- or period-specific parameters) to a reduced model (single set of parameters) suggested that parameter values of this relationship were different between the stands ($P < 0.001$), but not between the two periods within each stand ($P > 0.20$). Analysis of the least-squares problem of estimating the parameter values for the $Q_{10}$ function showed that increases in $R_b$ indeed must result in decreases in $a$ (see Appendix A).

To overcome the uncertainty in the estimates of $R_b$ and $a$ obtained from the temperature bins (pooled data, Fig. 3a), we used the more robust parameters obtained from the location data that covered the entire temperature range (Fig. 3b). The estimates of $F_H^a$ (with $T_S$ ranging from 5 to 23°C) calculated using the mean values of $R_b$ and $a$ from the location data deviated <2% from the estimates calculated using temperature-bin-specific parameterization. At PP, one set of parameters was sufficient to describe the $F_H^a$–temperature response (Table 2). At HW, the parameter values changed following the drought-breaking rains.
Table 2  Parameters of the responses of the forest floor CO₂ efflux (F₉) to soil temperature at 5 cm under nonlimiting soil moisture conditions F₉ₜₕ (Tₛ, Rₛ, a) and of the reduction function of soil moisture reduction function F₉ₜₕ to account for the effect of soil moisture F₉ₜₕ (θ, b, c) using daily mean θ.

<table>
<thead>
<tr>
<th>Period</th>
<th>Rₛ</th>
<th>a</th>
<th>Qₛ₀</th>
<th>b</th>
<th>c</th>
<th>RMSE</th>
<th>r²</th>
</tr>
</thead>
<tbody>
<tr>
<td>PP</td>
<td>1</td>
<td>0.841 ± (0.044)</td>
<td>0.103 ± (0.004)</td>
<td>2.80</td>
<td>32.04 ± (2.509)</td>
<td>3.64 ± (0.337)</td>
<td>0.490</td>
</tr>
<tr>
<td></td>
<td>2</td>
<td>0.902 ± (0.054)</td>
<td>0.102 ± (0.003)</td>
<td>2.77</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>HW</td>
<td>1</td>
<td>0.534 ± (0.026)</td>
<td>0.121 ± (0.003)</td>
<td>3.35</td>
<td>28.82 ± (3.239)</td>
<td>2.89 ± (0.453)</td>
<td>0.535</td>
</tr>
<tr>
<td></td>
<td>2</td>
<td>0.952 ± (0.062)</td>
<td>0.102 ± (0.004)</td>
<td>2.77</td>
<td></td>
<td></td>
<td></td>
</tr>
</tbody>
</table>

Parameter values for the F₉ₜₕ-Tₛ function are means over chamber locations. Rₛ is F₉ₜₕ at Tₛ = 0°C in μmol CO₂ m⁻² s⁻¹, where Tₛ is soil temperature at 5 cm. Parameters a, Qₛ₀ = e⁻ᵃθ⁻₀, b, and c are dimensionless. Residual mean squared error (RMSE) and r² refer to the model f(Tₛ,θ). Periods 1 and 2 refer to data before and after drought-breaking rains (08/27/02). For Rₛ and a, the standard error (SE, in parentheses) is that of the mean over the chamber positions (N = 12 at PP, and 30 at HW), and for b and c it is SE of the estimate from the nonlinear regression analysis (daily data, N > 50). Data are from the hardwood forest (HW) and the pine plantation (PP) located at Duke Forest. Capital letters refer to statistically significant differences between stands within a year, and lower case letters to between years within a stand; P < 0.05.

Effect of soil moisture on F₉ₜₕ

To assess the effect of low soil moisture on reducing the flux from its potential we analyzed the daily means of F₉ₜₕ because soil moisture measurements were available for the site but not for each chamber. Under nonlimiting soil moisture conditions changes in F₉ₜₕ were explained by an exponential function of temperature (r² > 0.96, filled symbols in Figs 4a and c). When soil moisture dropped to θ < 0.20 m³ m⁻³, F₉ₜₕ dropped below F₉ₜₕ (open symbols in Figs 4a and c). There was a more pronounced decrease in F₉ₜₕ with decreasing soil moisture at PP than at HW (insets, Figs 4a and c). The parameters (b and c in Eqn (1)) were statistically different between the sites (P < 0.001; Table 2).

In addition, frequently after major rain events (and when θ < 0.20 m³ m⁻³) the estimates of daily F₉ₜₕ were slightly smaller than the measured means for up to 5 days following the event (data not shown). Thus, the observed individual pulses in the measured F₉ₜₕ seemed related to rewetting of litter and soil. We modeled these patterns using Y = 1 + d₁ × D × e⁻ᵃθ₀, where Y is the ratio of measured F₉ₜₕ to F₉ₜₕ (Tₛ₉ₕ, θ₀), and D is number of days from the rain event (modified from Liu et al., 2002). The model explained reasonably well the residuals from values predicted based on Eqn (1) (r² > 0.80, P < 0.001).

Differences in monthly and annual F₉ₜₕ

The multiplicative model F₉ₕ (Tₛ₉ₕ, θ) with the mean values of Rₛ and a for the temperature response of F₉ₜₕ (with parameters representing pre- and postdrought-breaking rains), with the soil moisture reduction function, and with the correction for the effects of rain events, explained 92% and 93% of the variation in the measured daily mean F₉ₜₕ at PP and HW, respectively (Table 2, Figs 4b and d). There were no trends in the residuals of the model with respect to Tₛ₉ₕ, θ, or time (data not shown).

There were 304 days where F₉ₜₕ data were available concurrently at both stands. Over these days, the cumulative F₉ₜₕ was greater at HW (1490 g C m⁻²) than at PP (1140 g C m⁻²). The mean Tₛₕ was 17.2°C at HW, and 15.6°C at PP. The average daily maximum Tₛₕ was 0.5°C higher and maximum 2.3°C higher at HW compared with PP. Although the average value of θ was similar at PP (0.23 m³ m⁻³) and HW (0.22 m³ m⁻³), the frequency of low values of θ (< 0.20 m³ m⁻³) was higher at PP. At both stands, the modeled cumulative F₉ₜₕ during this period deviated < 1% from the measured value. Using the model, we found that of the difference of +269 g C m⁻² (Δ = HW-PP) -133 g C m⁻² was attributable to differences in the F₉ₜₕ-temperature response between the two stands, +260 g C m⁻² to the soil temperature difference, +120 g C m⁻² to the soil moisture response of F₉ₜₕ, and +76 g C m⁻² to soil moisture difference.

The monthly estimates of F₉ₜₕ suggested that F₉ₜₕ was higher at HW than at PP during the growing season in both years (Fig. 5). During the rest of the year, it was slightly higher at PP compared with HW or similar at the two stands. The intra-annual differences were roughly associated with differences in Tₛₕ between the stands, while the interannual variation resulted from different responses to drought. The annual estimates of F₉ₜₕ obtained using gap-filled data (Table 3) were similar to those based on only the model. Uncertainty around the estimates of F₉ₜₕ at each stand was obtained using the model results from each chamber location. Based on these, the annual F₉ₜₕ was not statistically different between years at each stand, or between stands in 2001, but was lower at PP than at HW in 2002 (P < 0.05).

We separated the effects of parameter differences on mean annual F₉ₜₕ between the stands from those caused...
Fig. 4  Daily mean forest floor CO$_2$ efflux ($F_{HF}$) as a function of soil temperature at 5 cm ($T_5$). The ratio of measured $F_{HF}$ to $F_{HF} (T_s)$ as a function of soil volumetric water content ($\theta$) in the insets (a, c). Measured $F_{HF}$ vs. modeled $F_{HF} (T_s, \theta)$ and a 1:1 line (b, d). Filled symbols represent data at $\theta \geq 0.20$ m$^3$ m$^{-3}$ and open symbols at $\theta < 0.20$ m$^3$ m$^{-3}$. Daily values averaged as in Fig. 1. Parameter values for the fitted functions are given in Table 2.

Fig. 5  Difference in the monthly sums of gap-filled forest floor CO$_2$ efflux ($F_{HF}$) and the monthly mean soil temperature ($T_5$, circles) between the hardwood forest (HW) and the pine plantation (PP).

by differences in the state variables (i.e., separating the effect of $R_b$, $a$, $b$, and $c$ from the effect of $\theta$ and $T_{soil}$) using the chain rule (see Appendix B). For 2001, the parameter difference ($-0.18 \mu$mol CO$_2$ s$^{-1}$ m$^{-2}$) was nearly entirely compensated by a $+0.17$ difference in the state. For 2002, the parameter difference ($-0.07$) was overwhelmed by the difference in state ($+0.16$). This analysis indicated that differences in $F_{HF}$ between the stands, estimated for average conditions using average parameters, were relatively small in 2001 and larger in 2002. This was a result of a decrease (from 2001 to 2002) in the relative difference between the parameter values of stands, thus, allowing the compensatory effect of the state variables to become dominating in 2002.

Discussion

Our study demonstrated that, under certain conditions, evergreen pine and deciduous broadleaf forest on similar soils and affected by the same climate, can emit different quantities of CO$_2$ from the forest floor. The
Table 3  Annual estimates of $F_H$ (g.C.m$^{-2}$.yr$^{-1}$) calculated using gap-filled data and standard deviation (in parentheses) obtained from the modeled estimates using each chamber location ($N = 12-30$).

<table>
<thead>
<tr>
<th>Year</th>
<th>$F_H$</th>
<th>$T_S$</th>
<th>$\theta$</th>
</tr>
</thead>
<tbody>
<tr>
<td>PP</td>
<td>2001</td>
<td>1336$^{+4}_{-5}$ (237)</td>
<td>13.9</td>
</tr>
<tr>
<td></td>
<td>2002</td>
<td>1231$^{+4}_{-5}$ (120)</td>
<td>14.3</td>
</tr>
<tr>
<td>HW</td>
<td>2001</td>
<td>1464$^{+4}_{-5}$ (230)</td>
<td>15.2</td>
</tr>
<tr>
<td></td>
<td>2002</td>
<td>1552$^{+4}_{-5}$ (254)</td>
<td>15.8</td>
</tr>
</tbody>
</table>

Data are from the hardwood forest (HW) and the pine plantation (PP) located at Duke Forest. $T_S$ is soil temperature ($^\circ$C) at 5 cm and $\theta$ is volumetric soil moisture (m$^3$.m$^{-3}$) measured over 0–30 cm at PP, and averaged over measurements at 10 and 25 cm at HW. Capital letters refer to statistically significant differences between stands within a year, and lower case letters to between years within a stand; $P < 0.05$.

The changes in $a$ with temperature are rooted in the sensitivity of the underlying respiration processes to temperature, but $F_H$ is de-coupled to some extent from belowground metabolism (e.g., Curiel Yuste et al., 2004). Thus, the temperature dependence of $a$ is probably a reflection of the form of the $Q_{10}$ function. An inherent mathematical property of the $Q_{10}$ function is that $a$ is related to $R_b$ (see Appendix A). Nevertheless, the simple $Q_{10}$ model performed equally well or better than the commonly used alternative, the modified Arrhenius model (Lloyd & Taylor, 1994) and was used in this study.

Except for conditions near soil saturation, $F_H$ is commonly found to be insensitive to soil moisture down to the moisture below which $F_H$ declines steeply as metabolic activity decreases (e.g., Hanson et al., 1993; Davidson et al., 1998; Fang & Moncrieff, 1999; Mielnick & Dugas, 2000). Although chamber-to-chamber variation in $\theta$ was not monitored in our study, it has been suggested that soil moisture is more important in explaining temporal variation than spatial variation in $F_H$ (Yim et al., 2003). Using the site mean $\theta$ for our stands explained reasonably well the reduction of $F_H$ from $F_H^0$ at $\theta < 0.20$ m$^3$.m$^{-3}$.

The increase in CO$_2$ efflux after rewetting is widely reported (e.g., Birch, 1958; Anderson, 1973; Borken et al., 1999; Borken et al., 2002; Lee et al., 2002; Rey et al., 2002; Curiel Yuste et al., 2003). The suggested causes include rapid evacuation of CO$_2$ from the soil pore space, increase in CO$_2$ production resulting from increase in substrate availability and microbial activity, and stimulation of root respiration. The effects of rewetting on annual soil CO$_2$ efflux in forests has, however, rarely been quantified. In a temperate spruce stand, rewetting after a simulated drought had no effect on the annual soil respiration during the first year, but increased it by 51% during the second year (Borken et al., 1999).

In this study, very low soil moisture might have affected $R_b$ and $a$ at both stands, but the effect was unclear and inconsistent, supporting the assumption of independent effects of soil temperature and moisture on $F_H$. On the other hand, the temperature sensitivity $a (= \ln(Q_{10})/10)$ of $F_H^0$, i.e., $F_H$ under nonlimiting soil moisture decreased with temperature at both stands, consistent with results from previous studies (Lloyd & Taylor, 1994; Qi et al., 2002; Janssens & Pilegaard, 2003).

$F_H^0$-temperature and $F_H$-moisture response

Because it was estimated from measurements of $F_H$, the temperature sensitivity of the efflux ($Q_{10}$) represents both CO$_2$ production and transport. Nevertheless, the $Q_{10}$ function describes $F_H$ reasonably well, even with parameters that do not vary with soil temperature, moisture, and season (e.g., Hanson et al., 1993; Davidson et al., 1998; Maier & Kress, 2000; Lee et al., 2002; Rey et al., 2002; Kang et al., 2003; Pumpen et al., 2003). This is surprising because an invariant base respiration ($R_b$) implies constant substrate pool size, and because $Q_{10}$ has been shown to decrease with temperature in plant tissues (e.g., Tjoelker et al., 2001) and $Q_{10}$ to depend on soil moisture (Borken et al., 1999; Qi et al., 2002; Reichstein et al., 2002a,b).

In this study, very low soil moisture might have affected $R_b$ and $a$ at both stands, but the effect was unclear and inconsistent, supporting the assumption of independent effects of soil temperature and moisture on $F_H$. On the other hand, the temperature sensitivity $a (= \ln(Q_{10})/10)$ of $F_H^0$, i.e., $F_H$ under nonlimiting soil moisture decreased with temperature at both stands, consistent with results from previous studies (Lloyd & Taylor, 1994; Qi et al., 2002; Janssens & Pilegaard, 2003).

Differences in $F_H$ between HW and PP

Higher mean $F_H$ in oak-dominated than pine-dominated stands has been linked to a greater annual $Q_{10}$ (Curiel Yuste et al., 2004). Curiel Yuste et al. (2004) demonstrated correlations between annual $Q_{10}$ and variables that reflect the seasonality of the stand, defined as the amplitude between the minimum and
maximum values of such variables. Among six sites varying in deciduousness, the amplitude of $F_{H}$ increased with the amplitude of LAI. Higher annual $Q_{10}$ in stands with greater LAI dynamics was assumed to reflect root phenology, implying that the variability in $Q_{10}$ among stands reflects differences in both temperature sensitivity and root phenology. Root phenology should be reflected to some extent in root biomass, presumably reaching maximum with LAI and resulting in high $R_b$ at the height of the growing season. Our analysis showed that variability in $R_b$ is predictably associated with variability in $Q_{10}$. Indeed, when data from the pine and oak stands were analyzed in 2 months intervals, $R_b$ was highest and $Q_{10}$ lowest in the middle of the growing season (Curriel Yuste et al., 2004).

Our results also showed that the inverse relationships between $R_b$ and $Q_{10}$ were only slightly different between the two stands. Consequently, if two stands show large difference in seasonality, and their $R_b$ and $Q_{10}$ are similar during the nongrowing season, the $R_b$ of the stand with greater seasonality should be higher and its $Q_{10}$ lower than the stand with lesser seasonality during the growing season. As predicted, $R_b$ at the oak stand studied by Curriel Yuste et al. (2004) tended to be higher than at the pine stand during the summer, but the temperature sensitivity was similar at both stands. Thus, the effect of root phenology on $F_{H}$ can be accounted for by analyzing $R_b$. However, $R_b$ can change for reasons unrelated to root phenology. In this study, $R_b$ at HW changed immediately following the drought-breaking rains in 2002, with no time lag necessary for increases in either LAI or root biomass. Our results suggested an alternative explanation to the intra-annual variability in $F_{H}$ that is also associated with the deciduousness of the forest, namely the effect LAI and thickness of the litter layer on soil temperature.

In the wetter of the 2 years, $F_{H}$ was statistically similar at the two stands. This was supported by the analytical evaluation of the controls over $F_{H}$ (Appendix B). The combined response to soil temperature and moisture at HW would have caused $F_{H}$ there to be lower compared with PP. However, this effect was compensated for by higher soil temperature and moisture at HW. Higher temperature throughout most of the year at HW was driven by differences in LAI dynamics and litter quality. While the peak LAI was somewhat higher at HW, the higher decomposition rate of broadleaf leaf litter (Finzi & Schlesinger, 2002) reduced the insulating litter layer at HW, bringing soil temperature in the uppermost layers closer to air temperature as has been shown elsewhere (Paul et al., 2004). In the nongrowing (leafless) season, a greater proportion of incoming radiation was reaching the ground resulting in a higher soil temperature at HW than that at PP, except when the radiative forcing was very low.

Although not significant, the difference in $F_{H}$ between the stands (10% in 2001) was consistent with the differences found between six broad-leaved forests and their adjacent coniferous counterparts, even though our data extended twofold the range of the published data (Raich & Tufekcioglu, 2000, Fig. 6a). (Adding the data from Curriel Yuste et al., 2004 did not change the overall relationship appreciably.) We evaluated the potential contributions of various C fluxes at each stand

![Fig. 6](image_url)
to this difference based on a simple mass balance approach. The sum of the litterfall C and allocation of C belowground, adjusted for changes in C stocks and the transport of C off site, should equal to the amount of C lost in $F_{H}$ (Raich & Nadelhoff, 1989; Giardina & Ryan, 2002). Annual $F_{H}$ and litterfall across forest types are positively correlated (Davidson et al., 2002; Fig. 6b). Although Davidson et al. (2002) found no correlation when mature temperate HW were analyzed separately, our HW data extended the range of the published data and, for the wetter of the 2 years, fell near the line of the overall correlation between $F_{H}$ and litterfall.

Assuming a steady state of C storage in the litter-soil system and averaging across biomes and age classes, Davidson et al. (2002) showed that belowground processes contribute the equivalence of ~75% of $F_{H}$ and the value is somewhat greater in young than mature stands. Our data suggested that in 2001, these processes accounted for ~76% of $F_{H}$ at both stands. Because we observed little leaf litter on the soil surface at HW just prior to leaf abscission in the autumn, we assumed that C is not accumulating in the litter-soil system, and the difference between $F_{H}$ and litterfall (1122 g m$^{-2}$ yr$^{-1}$ in 2001) represented C allocation belowground and belowground contribution to $F_{H}$ at this stand. In contrast, it is certain that PP was accumulating C in the litter-soil system. Averaged over the last few years, the PP accumulated C in both the litter layer (44 g m$^{-2}$ yr$^{-1}$; Schlesinger & Lichter, 2001) and the upper 30 cm of the mineral soil (138 g m$^{-2}$ yr$^{-1}$; Lichter et al., in press). Using these annual C accumulation rates together with litterfall and $F_{H}$ data for 2001, the mass balance showed that C allocation belowground at PP was ~1180 g m$^{-2}$ yr$^{-1}$. Thus, in 2001, C input belowground was somewhat higher at PP, and the lower $F_{H}$ at PP was related to it being in the C accumulation stage, in contrast to HW that reached a steady state of C storage.

During the drought year of 2002, the large difference in $F_{H}$ between the stands (26%) resulted from a relatively smaller difference in the response of $F_{H}$ to soil temperature and moisture and a larger effect of the difference in those state variables (see Appendix B). The difference in the response of $F_{H}$ to soil temperature decreased following the drought-breaking rains in August 2002. The value of $F_{H}$ at HW increased to a value similar to that at PP, indicating an increase in substrate availability related to cessation of the drought. Results of girdling studies, terminating the supply of recently assimilated C to roots, have suggested that recent C contributes a significant amount to $F_{H}$ (Hogberg et al., 2001; Bhupinderpal-Singh et al., 2003). Previous studies at our stands have shown that the canopy conductance at PP is more sensitive to soil moisture than that at HW (Oren et al., 1998; Pataki & Oren, 2003). This may have resulted in smaller impact of the drought on C assimilation at HW, and continued allocation of recent C to belowground where it accumulated in the soil to a greater extent than at PP. In summary, the similar responses to soil temperature and moisture at the two stands in the latter half of 2002 allowed the differences in soil temperature and moisture to yield much higher $F_{H}$ at HW compared with PP.

We repeated the mass balance calculation to assess the contribution of belowground processes to $F_{H}$ in 2002. At PP, we assumed that the decomposition of litter contributed to $F_{H}$ the same amount of C as it did in the wetter year of 2001 (~288 g C m$^{-2}$ yr$^{-1}$), 38 g C m$^{-2}$ yr$^{-1}$ in excess of the litterfall in 2002. This would eliminate C accumulation in the litter layer maintaining the C storage there unchanged (998 g m$^{-2}$; Lichter et al., in press). If annual C accumulation in the mineral soil remained the same as the annual average, allocation of C belowground ($= F_{H}$-litter decomposition + accumulation) would have been ~1081 g m$^{-2}$ yr$^{-1}$. If the greater sensitivity of canopy conductance to drought at PP (Oren et al., 1998; Pataki & Oren, 2003) translates to a reduced belowground allocation, annual C accumulation in the mineral soil would have to decrease accordingly. Even though HW might be less sensitive to drought than PP, it is unlikely that it would be able to allocate belowground larger quantities of C in 2002 than in 2001 (~1263 vs. ~1122 g m$^{-2}$ yr$^{-1}$, respectively). A better assumption may be that allocation was similar to that in 2001. If so, this represents a loss of ~141 g C m$^{-2}$ yr$^{-1}$ from the mineral soil of HW in 2002, more than the estimated loss from the litter layer at PP (~38 g C m$^{-2}$ yr$^{-1}$).

Conclusions

The aim of this study was to assess the role of vegetation in affecting seasonal and annual $F_{H}$ in a maturing loblolly PP and a mature Oak-Hickory type forest. The results showed that with the exception of the winter months, $F_{H}$ was higher at HW compared with the adjacent PP. The higher $F_{H}$ at HW resulted from the combined effects of stand differences in the responses to soil temperature and moisture, and the frequency distribution of soil temperature and moisture. The cumulative effect of the different responses was reflected in the interannual variation in $F_{H}$ at the two stands. Our results showed that predictions of $F_{H}$ under climate change scenarios, including changes in soil temperature and drought severity, are complicated by the differential effect of drought on litter production and belowground C allocation, and by the likelihood

© 2005 Blackwell Publishing Ltd, Global Change Biology, 11, 421-434
that a series of drought years would generate responses that are quite different from the response of one severe drought year among normal years. Although these results indicate that C accumulation in litter and soil of PPs in this region is likely to be higher than in mature HW, it is essential to consider the fate of soil C already stored in broad-leaved forests as these undergo large-scale conversion to pine plantations.

Acknowledgements

This research was supported by the Office of Science (BER), US Department of Energy, through the Terrestrial Carbon Processes Program (TCP) Grant no. DE-FG02-01ER63015 and Grant no. DE-FG02-95ER62083, and through BER's Southeast Regional Center (SERC) of the National Institute for Global Environmental Change (NIEEC) under Cooperative Agreement no. DE-FG02-03ER63613, and by the US Forest Service through the Southern Global Change Research Program and the Southern Research Station. The work of the senior author has been supported by grant 52228 from the Academy of Finland.

References


SOIL CO₂ EFFLUX IN PINE AND HARDWOOD FORESTS 433


Reichstein M, Tenhunen JD, Roupsard O et al. (2002b) Severe drought effects on ecosystem CO₂ and H₂O fluxes at three Mediterranean evergreen sites: revision of current hypotheses? Global Change Biology, 8, 999–1017.


Appendix A: Theoretical relationship between a and b

Given an independent variable x and a dependent variable y, the least-squares problem of determining the slope (m) and intercept (b) of y = mx + b from n (>3) observations of x and y (hereafter referred to as xᵢ, yᵢ)

© 2005 Blackwell Publishing Ltd, Global Change Biology, 11, 421–434
can be formulated as the system of two equations with two unknowns:

\[ \frac{1}{n} \sum_{i=1}^{n} y_i = m \frac{1}{n} \sum_{i=1}^{n} x_i + b, \]  
(A.1)

\[ \frac{1}{n} \sum_{i=1}^{n} x_i y_i = m \frac{1}{n} \sum_{i=1}^{n} x_i^2 + b \frac{1}{n} \sum_{i=1}^{n} x_i, \]  
(A.2)

Letting \( \Psi = \frac{1}{n} \sum_{i=1}^{n} \Psi_i \), where \( \Psi \) is either \( x \) and \( y \) and adopting this notation in Eqn (A.1) yields a relationship between \( b \) and \( m \) as

\[ b = \bar{y} - m \bar{x}. \]  
(A.3)

The parameter estimation for \( R_b \) and \( a \) in Eqn (1) with no soil moisture limitations (or at specified moisture or temperature bins) can be formulated as a linear regression problem if \( b = \ln(R_b), y = \ln(F_H), m = a \), and \( x = T_{soil} \). Replacing these equalities into Eqn (A.3) yields

\[ R_b = e^{[\ln(\bar{y}) - m \ln(\bar{x})]}, \]  
(A.4)

Note, \( \ln(\bar{y}) \neq \ln(\bar{y}) \). Because \( F_H \) ranges between 2 and 4 \( \mu \text{mol CO}_2 \text{m}^{-2} \text{s}^{-1} \) for most of the times with maximum excursions to about 12 \( \mu \text{mol CO}_2 \text{m}^{-2} \text{s}^{-1} \), the variability in \( \ln(F_H) \) becomes much smaller (i.e., at least one order of magnitude) than the variability in \( F_H \). Hence, to a first approximation \( e^{\ln(F_H)} \) becomes approximately a constant, \( \gamma \). Upon replacing in Eqn (A.4), the relationship between \( R_b \) and \( a \) becomes

\[ R_b = \gamma e^{-a T_{soil}}. \]  
(A.5)

It is evident from the above model that increases in \( a \) must result in decreases in \( R_b \). Eqn (A.5) can be inverted to obtain \( a \) as a function of \( R_b \):

\[ a = \frac{1}{T_{soil}} \ln \left( \frac{\gamma}{R_b} \right). \]  
(A.6)

In Fig. 7, the values of \( a \) calculated based on Eqn (A.6) \((a')\) are shown to be in close agreement with those obtained from the curve fitting of data in Fig 3b.

Appendix B: Parameter vs. state effects on \( F_H \)

When comparing differences in \( F_H \) between HW and PP, the genesis of the potential differences can be attributed to two types; state differences (i.e., the state variables \( \theta \) and \( T_{soil} \)) and parameter differences (i.e., \( R_b, a, b, \) and \( c \)). To separate these two effects, we start with Eqn (1) and proceed to quantify the variability in \( F_H \) as state and/or parameter using the chain rule

\[ dF_H = \left[ \frac{\partial F_H}{\partial R_b} \right] dR_b + \left[ \frac{\partial F_H}{\partial a} \right] da + \left[ \frac{\partial F_H}{\partial b} \right] db + \left[ \frac{\partial F_H}{\partial c} \right] dc \]  
(B.1)

\[ \text{Parameter change} \]

\[ + \left[ \frac{\partial F_H}{\partial T_{soil}} \right] dT_{soil} + \left[ \frac{\partial F_H}{\partial \theta} \right] d\theta, \]

\[ \text{State change} \]

where

\[ \frac{\partial F_H}{\partial R_b} \bigg|_{a,b,c,T_{soil},\theta} = e^{\gamma T_{soil}} \left[ 1 - e^{-(b+c)} \right]. \]  
(B.2)

\[ \frac{\partial F_H}{\partial a} \bigg|_{R_b,b,c,T_{soil}} = R_b T_{soil} e^{\gamma T_{soil}} \left[ 1 - e^{-(b+c)} \right]. \]  
(B.3)

\[ \frac{\partial F_H}{\partial b} \bigg|_{R_a,a,c,T_{soil}} = \theta R_b e^{\gamma T_{soil}} \left[ e^{-(b+c)} \right]. \]  
(B.4)

\[ \frac{\partial F_H}{\partial c} \bigg|_{R_b,a,c,T_{soil}} = R_b e^{\gamma T_{soil}} \left[ e^{-(b+c)} \right]. \]  
(B.5)

\[ \frac{\partial F_H}{\partial T_{soil}} \bigg|_{R_b,a,b,c} = R_b \theta e^{\gamma T_{soil}} \left[ 1 - e^{-(b+c)} \right]. \]  
(B.6)

\[ \frac{\partial F_H}{\partial \theta} \bigg|_{R_b,a,b,c} = b R_b e^{\gamma T_{soil}} \left[ 1 - e^{-(b+c)} \right]. \]  
(B.7)

With these estimates and assuming that \( d \Psi \approx \Delta \Psi \) (i.e., a first-order linear Taylor series expansion), where \( \Psi \) is any of the five variables \( \theta, T_{soil}, R_b, a, b, \) and \( c \) the relative change is given by

\[ \frac{\Delta F_H}{F_H} = \left[ \frac{\Delta R_b}{R_b} + T_{soil} \Delta a + \frac{\Delta b}{b} \Delta \theta \right] \frac{d \Psi}{\Psi} + \frac{\Delta c}{c} \frac{\Delta \Psi}{\Psi}, \]  
(B.8)

\[ \text{Parameter change} \]

\[ + \left[ \frac{\Delta T_{soil}}{T_{soil}} \right] \frac{d \Psi}{\Psi} + \left[ \frac{\Delta \theta}{\theta} \right] \frac{d \Psi}{\Psi}, \]

\[ \text{State change} \]

where the annual values of \( \theta, T_{soil}, R_b, a, b, \) and \( c \), averaged for the two stands, represented a ‘reference stand’, and \( \Delta \) referred to the differences between HW and PP \((\Delta = \text{HW} - \text{PP})\).